

Optical properties of atmospheric aerosols in León (Spain)



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INTRODUCTION

Nowadays, there is a clear evidence that aerosols directly and indirectly impact on thermodynamic processes and radiative fluxes of the Earth's atmosphere (Andreae & Rosenfeld, 2008). Aerosols influence atmospheric radiative balance (Fig. 1) directly through the interaction of radiation and particulate matter, such as scattering and absorption, and indirectly some act as cloud condensation and ice nuclei. Aerosol properties depend on the size of the particles and some of them, like the optical properties,

		RADIATIVE FORCING COMPONENT	s		
	RF Terms		RF values (W m ⁻²)	Spatial scale	LOSU
	Long-lived greenhouse gases	CO₂ ⊫–1	1.66 [1.49 to 1.83]	Global	High
		N ₂ O CH ₄ Halocarbons	0.48 [0.43 to 0.53] 0.16 [0.14 to 0.18] 0.34 [0.31 to 0.37]	Global	High
lic	Ozone	Stratospheric	-0.05 [-0.15 to 0.05] 0.35 [0.25 to 0.65]	Continental to global	Med
Anthropogenic	Stratospheric water vapour from CH ₄	H	0.07 [0.02 to 0.12]	Global	Low
Anthro	Surface albedo	Land use H Black carbon on snow	-0.2 [-0.4 to 0.0] 0.1 [0.0 to 0.2]	Local to continental	Med - Low
	Total		-0.5 [-0.9 to -0.1]	Continental to global	Med - Low
	Aerosol Cloud albedo effect		-0.7 [-1.8 to -0.3]	Continental to global	Low
	Linear contrails		0.01 [0.003 to 0.03]	Continental	Low
Natural	Solar irradiance		0.12 [0.06 to 0.30]	Global	Low
	Total net anthropogenic		1.6 [0.6 to 2.4]		
	-2	-1 0 1 2			

EXPERIMENTAL



1. Determination of the hours with a **solar** elevation angle above **10°.**

2. Particle size distributions have been obtained with an optical spectrometer PCASP-X. Diameters

Fine mode **Coarse mode Total**

3. Definition and characterization of aerosol size modes during the hours with a solar elevation angle



July 2016	1.545-0.006i	2.029
August 2016	1.547-0.006i	2.053
September 2016	1.552-0.008i	2.154
October 2016	1.560-0.013i	1.821
November 2016	1.556-0.012i	1.842
December 2016	1.559-0.013i	1.792
January 2017	1.550-0.009i	1.952
February 2017	1.544-0.007i	2.023

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Figure. 4. Aerosol size distribution of particles in

8.00

Month/Year [–]			
Month, Icai	StationE1 (TRAFFIC)	Station E4 (BACKGROUND)	Increase (%)
March/16	17	11	55
April/16	13	7	86
May/16	14	10	40
June/16	17	11	55
July/16	20	15	33
August/16	20	15	33
September/16	23	17	35
October/16	22	11	100
November/16	24	11	118
December/16	31	16	94
January/17	30	15	100
February/17	24	12	100
Annual	21	13	62

- The **density** in the summer months increases with a **maximum** in September of 2.154 g cm⁻³.
- The **imaginary** part of the refractive index in summer months varies between 0.006 and **0.008**, that means that the aerosols present have

the capacity to absorb the radiation (Table 1).

The highest values for MSE, MAE, MEE and MBSE were reached in winter for 440 nm, with values of 3.3±2.3, 0.3±0.1, 3.5 ± 2.4 and 0.1 ± 0.1 Mm⁻¹, respectively (Fig. 5).

spring, summer, autumn and winter.

Saharan dust intrusions increase the number of particles, promoting higher values of the optical parameters in summer (Alonso-Blanco et al., 2018) (Fig. 5).

The increase in traffic, the use of heating devices and the reduction in the mixing layer thickness promote high PM_{10} concentrations in the winter months (Table 2).

Conclusions

The density of the particles in León was higher in the summer months with values higher than 2 g cm⁻³. In summer, aerosol particles presented higher capacity to absorb the a radiation. The highest values for MSE, MAE, MEE and MBSE were reached in winter for 440 nm. These results, together with those obtained from the ultrafine mode, will be implemented in the Global Atmospheric Model (GAME) to analyze the effect the evolution of aerosol particles on the radiative forcing.

Domestic heater devices

High traffic intensity

Lower altitude of mixing layer thickness

 $PM_{10} (\mu g m^{-3})$



Figure 5. Mass efficiencies of MSE, MAE, MEE and MBSE as a function of three different wavelengths (440 nm, 670 nm, 870 nm respectively) for summer, autumn, winter and spring.

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